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MAXDOAS formaldehyde slant column measurements during CINDI: intercomparison and analysis improvement

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Abstract

We present intercomparison results for formaldehyde (HCHO) slant column measurements performed during the Cabauw Intercomparison Campaign of Nitrogen Dioxide measuring Instruments (CINDI) that took place in Cabauw, the Netherlands, in summer 2009. During two months, nine atmospheric research groups simultaneously operated MAXDOAS instruments of various designs to record UV-visible spectra of scattered sunlight at different elevation angles that were analysed using common retrieval settings. The resulting HCHO dataset was found to be highly consistent, the mean difference between instruments generally not exceeding 15 % or 7.5×10^{15} molec cm^{-2} , for all viewing elevation angles. Furthermore, a sensitivity analysis was performed to investigate the uncertainties in the HCHO slant column retrieval when varying key input parameters such as the molecular absorption cross-sections, correction terms for the Ring effect or the width and position of the fitting interval. This study led to the identification of potentially important sources of errors associated with cross-correlation effects involving the Ring effect, O_4 , HCHO and BrO cross-sections and the DOAS closure polynomial. As a result, a set of updated recommendations was formulated for HCHO slant column retrieval in the 336.5–359 nm wavelength range. To conclude, an error budget is proposed which distinguishes between systematic and random uncertainties. The total systematic error is estimated to be of the order of 20 % and is dominated by uncertainties in absorption cross-sections and related spectral cross-correlation effects. For a typical integration time of one minute, random uncertainties range between 5 % and 30 %, depending on the noise level of individual instruments.

1 Introduction

Formaldehyde (HCHO) is the most abundant organic carbonyl compound in the atmosphere. Being a short-lived oxidation product of a large number of Volatile Organic Compounds (VOCs), its abundance can be closely related to VOC emissions of

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natural origin or from human activities. In particular, the variability of HCHO over continents is dominated by local production from non-methane volatile organic compounds (NMVOCs) (Millet et al., 2006; Stavrou et al., 2009a,b and references therein), originating from biogenic, pyrogenic or anthropogenic sources. These emissions, through their involvement in tropospheric ozone chemistry, are important in controlling air quality. Over Europe, satellite observations reveal a marked seasonal cycle, with a summer maximum and a winter minimum (Curci et al., 2010).

Tropospheric formaldehyde has been observed from nadir UV backscatter satellite sensors (Chance et al., 2000; Palmer et al., 2003; Wittrock et al., 2006; De Smedt et al., 2008, 2010; Millet et al., 2008; Marbach et al., 2009; Vrekoussis et al., 2010; Curci et al., 2010), as well as from airborne and ground-based in-situ instruments (e.g. Hak et al., 2005; Fried et al., 2011 and references therein) and ground-based remote-sensing Fourier transform infrared spectrometers (e.g. Demoulin et al., 1999; Jones et al., 2009; Vigouroux et al., 2009). In the UV, HCHO can be measured using active and passive differential optical absorption spectroscopy (DOAS) (see Platt and Stutz, 2008, for a review). Among the category of passive DOAS systems, the Multi-AXis DOAS (MAXDOAS) technique can be used for tropospheric trace gas monitoring (Van Roozendaal 2003; Hönninger et al., 2004; Sinreich et al., 2005). The MAXDOAS technique has been successfully applied to NO₂, HCHO, SO₂, BrO, and CHOCHO detection (e.g. Bobrowski et al., 2003; Wittrock et al., 2004; Heckel et al., 2005; Sinreich et al., 2007; Pikel'naya et al., 2007; Theys et al., 2007; Clémer et al., 2009; Vlemmix et al., 2010, 2011a; Irie et al., 2011; Wagner et al., 2011).

MAXDOAS instruments perform quasi-simultaneous observations of scattered sunlight in a range of viewing directions scanned from the horizon to the zenith. Measurements made at low elevation angles have a higher sensitivity to lower tropospheric layers, since in this geometry the photons travel longer paths through these layers than photons sampled at larger elevation angles. This results in an increased sensitivity to atmospheric absorbers that are located near the surface, such as HCHO. It has been shown that the MAXDOAS technique can be used to retrieve information on

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the vertical distribution of both trace gases and aerosols (Wagner et al., 2004; Frieß et al., 2006; Irie et al., 2008, 2009; Clémer et al., 2010; Li et al., 2010). These advantages, combined with the fact that measurements can be made automatically year-round with a good/frequent temporal sampling, make the MAXDOAS method suitable for long-term monitoring of atmospheric composition. Moreover, due to their spatial representativeness in both vertical and horizontal axes (Irie et al., 2011; Wagner et al., 2011), MAXDOAS measurements are also particularly well suited to make the link between satellite observations, chemical transport model calculations, and in-situ measurements in a validation context (Brinksma et al., 2008; Vlemmix et al., 2011b).

The first reported MAXDOAS measurements of HCHO were performed in the Italian Po-Valley in summer 2002 and 2003 (Heckel et al., 2005; Wagner et al., 2011). Soon after, Hak et al. (2005) reported on an intercomparison exercise including the chromatographic technique, long-path (LP) DOAS, Fourier Transform Infrared (FTIR), MAXDOAS, and Hantzsch-type in-situ instruments, showing good consistency between the different techniques. In 2006, Wittrock et al. published a study on global HCHO mapping from the SCIAMACHY nadir satellite instrument including comparisons with two MAXDOAS systems (in Kenya and the Netherlands), and found good agreement within the estimated uncertainties. Pikel'naya et al. (2007) and Inomata et al. (2008) also compared MAXDOAS HCHO with LP-DOAS and PTR-MS data, respectively, and Vigouroux et al. (2009) compared SCIAMACHY, MAXDOAS and FTIR HCHO total columns in a tropical remote region. Recently, Fried et al. (2011) presented a comparison of ground-based MAXDOAS and aircraft HCHO measurements during the INTEX-B campaign and showed good agreement between both datasets.

Although the aforementioned studies have largely demonstrated the capabilities of MAXDOAS systems for HCHO detection, no effort has been devoted so far to intercompare and harmonize instruments and retrieval methods. Such activities are an essential prerequisite for the reliable retrieval of vertical columns and profiles from MAXDOAS HCHO measurements. The Cabauw Intercomparison Campaign of Nitrogen Dioxide measuring Instruments (CINDI) in the Netherlands in June–July 2009 (Roscoe et al.,

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2010; Piters et al., 2012), where a number of MAXDOAS instruments were jointly operated, presented the opportunity to assess for the first time the consistency of these HCHO measuring systems. The Cabauw site is located in a semi-rural area of the Netherlands where formaldehyde concentrations are expected to be between one and several tens of ppbv, which are typical background levels in the continental boundary layer and urban regions, respectively (Hak et al., 2005; Fried et al., 2011). In their study on multi-component MAXDOAS retrievals during CINDI, Irie et al. (2011) reported median HCHO Vertical Mixing Ratios (VMRs) of around 2.5 ppbv, and peak values of up to 8 ppbv. The CINDI data set is therefore ideally suited for an HCHO intercomparison exercise.

This paper is structured as follows. Section 2 gives an overview of the CINDI campaign, the data recorded, the instruments, and the settings used by the different groups. Results from the HCHO slant column intercomparison are then presented in Sect. 3. In Section 4, we focus on sensitivity tests performed to assess the stability of the HCHO evaluations. Results are used to derive an error budget for slant column retrievals. Conclusions and perspectives are given in Sect. 5.

2 Data analysis

2.1 The CINDI intercomparison campaign

The CINDI campaign took place at the Cabauw Experimental Site for Atmospheric Research (CESAR site; Russchenberg et al., 2005) a semi-rural area in the Netherlands. Although the campaign mainly focused on tropospheric NO₂, aerosols were also monitored with in-situ systems, sun photometers and a lidar (Piters et al., 2012; Zieger et al., 2011; Irie et al., 2011). To complement the atmospheric composition measurements, the site offered a large ensemble of meteorological observations. A detailed description of the CINDI campaign and an overview of first results can be found in Piters et al. (2012).

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Instruments were installed and tested between 8 and 14 June 2009 and a formal semi-blind intercomparison of NO₂ and O₄ Differential Slant Column Densities (DSCDs) took place from 15 to 30 June 2009 (Roscoe et al., 2010). After that date, additional measurements were performed by several groups until 24 July, focusing on other species, such as HCHO, BrO, and CHOCHO. The goal of the formal semi-blind NO₂ and O₄ intercomparison was to characterize the current level of consistency of multi-axis instruments, as previously done for the zenith-sky instruments of the Network for the Detection of Atmospheric Composition Change (NDACC; e.g. Roscoe et al., 1999; Vandaele et al., 2005). During CINDI, 22 NO₂ measuring systems were intercompared, and most of them were shown to meet the criteria for endorsement by NDACC (Roscoe et al., 2010).

2.2 Instruments

The present work focuses on HCHO measurements from MAXDOAS instruments. Nine groups participated in this exercise, as reported in Table 1. In addition NIWA (National Institute for Water and Atmospheric Research, New Zealand) and KNMI (Koninklijk Nederlands Meteorologisch Instituut, the Netherlands) performed a few HCHO measurements during the campaign. However, these data were too sparse to include in the intercomparison.

A comprehensive description of the CINDI MAXDOAS instruments can be found in Piters et al. (2012) and Roscoe et al. (2010), and additional references relevant for the present study are given in Table 1. All systems that operated during CINDI recorded spectra at a set of prescribed elevation angles (2°, 4°, 8°, 15°, 30° and the zenith), and at a fixed azimuth angle of 287° relative to North. A full cycle of MAXDOAS measurements was generally obtained within half an hour. For the intercomparison of HCHO, only measurements recorded at solar zenith angle (SZA) less than 75° were used in order to exclude error-prone twilight measurements that are not relevant for the present study.

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As can be seen in Table 1, the characteristics of the various instruments were quite diverse, ranging from commercial mini-DOAS systems to state-of-the-art scientific grade instruments equipped with thermoelectrically cooled CCD array detectors of large dimension. Accordingly, the stability and noise levels were expected to differ considerably from one system to another, with a direct impact on corresponding HCHO measurements. Another important point is that some instruments were set up to record spectra in both UV and visible ranges simultaneously, while others had to alternate between these two wavelengths regions. As a result, some systems (e.g. INTA and Toronto) could only measure HCHO after 30 June, at the end of the formal period of the NO₂ intercomparison exercise described in Roscoe et al. (2010). Others instruments (WSU, NASA) switched their wavelength range repeatedly over time, alternating the two types of measurements throughout the campaign. It should be noted that the MAXDOAS instrument of the Toronto group was tested in multi-axis mode for the first time during this campaign, and only a few days of data were available.

2.3 HCHO analysis settings

Formaldehyde presents structured absorption bands in the UV spectral region, between 240 and 360 nm. Because of the increased ozone absorption below 320 nm, only the bands above 320 nm are generally used for the spectral retrieval of HCHO. In this region, the HCHO DSCDs are retrieved from measured spectra by applying the DOAS technique (Platt and Stutz, 2008, and references therein). The wavelength interval for fitting is generally optimized so as to (1) maximize the sensitivity to HCHO, (2) minimize the fitting residuals and the scatter of the retrieved HCHO slant columns, and (3) minimize the interferences with other absorbers. Figure 1 displays the absorption cross-sections of the trace gases that need to be considered in the 320–370 nm wavelength region. Based on experience and heritage from past studies on ground-based (e.g. Heckel et al., 2005; Vigouroux et al., 2009) and satellite instruments (Wittrock et al., 2006; De Smedt et al., 2008, 2009), standardized analysis settings were defined and prescribed for use in the present intercomparison exercise. These baseline settings

are summarized in Table 2. Note that all absorption cross-sections were convolved to the resolution of the individual instruments using slit function information provided by each group. In the case of the O₃ and NO₂ cross-sections, a solar I₀ correction was applied (Aliwell et al., 2002).

For the retrievals, daily reference spectra were taken from the zenith observations around 11:40 UT. Figure 2 presents an example of HCHO fitting results obtained with the BIRA-IASB instrument on 30 June 2009, at 4° elevation angle and 43° SZA. The corresponding residuals (approximately 10⁻⁴ RMS) are typical of low-noise scientific grade instruments. Under similar conditions, residuals can be an order of magnitude larger when using compact mini-DOAS systems.

2.4 Results

Figure 3a shows an example of HCHO measurements obtained during CINDI. HCHO DSCDs measured by the BIRA-IASB MAXDOAS instrument at 30° elevation were converted into equivalent vertical columns (VCDs) using a simple geometrical approximation (Hönninger et al., 2004). The time-series of HCHO VCDs covers approximately one month during summer. Comparing the retrieved columns with measurements of the ambient temperature (Fig. 3a), one can see that the HCHO variations are strongly correlated with corresponding changes in the temperature. This feature was reported by Irie et al. (2011) and is consistent with a background HCHO source, mainly due to the oxidation of biogenic NMVOCs. In contrast, the HCHO peak values as observed on several days in July are probably related to anthropogenic emission sources. In Fig. 3b, the complete VCD dataset was binned and plotted as a function of the local time. As can be seen from the standard deviations, the natural variability of HCHO is large, but the diurnal variation consistently peaks in the afternoon, which is likely related to the progressive oxidation of the NMVOCs emitted during the day.

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3 Slant column intercomparison results

We follow the method introduced in Roscoe et al. (2010) for the NO₂ and O₄ slant column measurement intercomparison. HCHO DSCDs retrieved by each group using the baseline analysis settings defined in Table 2, are averaged over periods of 30 min.

This procedure minimizes the impact of the temporal and spatial variability of HCHO and of the differences in read-out noise between instruments. An example of the diurnal evolution of the resampled HCHO DSCDs is shown in Fig. 4 for different elevation angles and for SZA < 75°, on 2 July 2009. This day was chosen because almost all instruments were in operation and the HCHO concentrations had a smooth variation over time. As can be seen, the HCHO DSCDs are consistently larger at low elevation due to the enhanced light path in the near-surface HCHO layer; good agreement is observed between various datasets.

To proceed further, the measurements from each group were first compared to the Bremen dataset, which has the most complete time-series, very frequent measurements and smooth behaviour. In the second step, a reference dataset was created by grouping the instruments that presented the best mutual agreement. To illustrate this procedure, Fig. 5 displays the BIRA HCHO DCSDs plotted against those of Bremen, for different elevation angles. Relevant statistical parameters (correlation coefficient, slope and intercept of the linear regression fit) are indicated. As can be seen, the scatter plots are compact, and the slopes and correlation coefficients are close to unity, confirming the good agreement between the Bremen and BIRA instruments. Statistical parameters at 30° and 90° elevation are slightly smaller, due to the smaller dynamical range of the DSCDs at those elevations (see Fig. 4). Similar agreement was found with INTA, and therefore a reference data set was created by averaging data from the Bremen, BIRA and INTA instruments. The scatter plots displayed in Fig. 6 illustrate, for the 4° elevation case, the results of the final comparison where data from each individual instrument are compared to the merged reference. Although the number of coincident points can differ greatly depending on instruments, the data of most of the

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groups agree quite well with the reference. Figure 7 presents the final results of the statistical analysis, constructed using the whole ensemble of off-axis measurements (2° , 4° , 8° , 15° and 30°). Most instruments compare relatively well with the reference for most of the elevation angles: correlation coefficients are close to unity (illustrating the compactness of the scatter plot with respect to the reference), slopes deviate by no more than 15% from the reference, and intercepts are close to zero. Note that larger relative differences against the reference are systematically obtained at 30° elevations due to the lower HCHO DSCDs at this elevation angle (see Fig. 4). Also smaller values for the correlation coefficients and larger intercept values and corresponding uncertainties are found for the JAMSTEC, NASA and Toronto datasets, which might possibly be connected to the larger noise of the corresponding instruments with respect to others (see Fig. 6). A test with an orthogonal regression (instead of linear) shows similar results.

For each instrument, the histograms of the absolute HCHO DSCDs deviations are presented in Fig. 8. Only the results at 4° elevation are shown, but similar conclusions can be drawn at other elevations. All instruments, except the Toronto one, have a symmetric and quasi-Gaussian shape for the 4° elevation, but large differences in the Gaussian FWHM (full width at half maximum). Largest FWHM values are found for JAMSTEC, NASA and Mainz, consistent with the noise level of those instruments. Some of the groups (IUPH, JAMSTEC, NASA, Toronto and Mainz) also display a small but significant bias compared to the reference. It should be mentioned again that the first three instruments (BIRA, INTA and Bremen) were used to construct the reference dataset and thus a better agreement with the reference is expected due to this choice. For the Mainz instrument, the negative bias might be related to the position of the instrument, at an altitude of 20 m on the Cabauw tower, where measurements probably loose sensitivity to HCHO present at the surface. In case of the Toronto instrument, the statistical results should be interpreted with care since there were only five days of measurements.

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The HCHO intercomparison exercise presented here should be connected to the NO_2 and O_4 comparisons performed by Roscoe et al. (2010). They reported fitted slopes for the visible MAX-DOAS instruments within 10 % of unity at almost all non-zenith elevations, with most instruments within 5 %; similarly, values of 12 % and 7 %, were found for the UV sensors. The 15 % deviation from unity slope derived in our study for the HCHO columns in the UV are thus a good result, considering that retrieving HCHO DSCDs is more challenging than retrieving NO_2 , because of the smaller optical thickness of HCHO.

4 Sensitivity study

The intercomparison study presented in Sect. 3 has been conducted using standardized retrieval settings consistently applied by each participating group. This harmonization procedure allowed us to concentrate on instrument-specific differences and to draw conclusions on the overall level of consistency between HCHO measuring systems operated during CINDI. We now concentrate on evaluating the sensitivity of HCHO results to possible changes in the retrieval settings using representative spectra from the BIRA instrument recorded on a clear day (4 July 2009). As will be illustrated hereafter, these sensitivity tests highlight possible optimizations in the HCHO slant column retrieval parameters and lead to the recommendation of new analysis settings. Sensitivity test results are shown both on DSCDs and on dDSCDs, i.e. the difference of the off-axis DSCDs with the zenith DSCDs of the scan (the closest in time), as it's the information used in the profile inversions.

4.1 Degree of closure polynomial and Ring effect

When performing a DOAS retrieval, an important free parameter is the degree of the polynomial function that is used to account for the smooth part of the attenuation spectrum. To avoid oscillations that may correlate with trace gas absorption features, the

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degree of this polynomial is generally restricted to values less than 5. For the baseline retrievals, a 3rd order polynomial was selected. However, during our sensitivity tests, we noticed that any changes to these polynomial settings had a strong impact on the diurnal behaviour of the HCHO DSCD, especially for high elevation angles including zenith, as can be seen in Fig. 9a, b, c. This raised the following two questions: (1) why such a dependence on the polynomial order, and (2) which one of the tested settings is the most satisfactory? In order to answer the second question, we decided to investigate the consistency of VCDs estimates as follows. In a first approximation, the HCHO VCD can be derived from measured DSCDs in two different simple ways: first, from the difference between 30° elevation off-axis and zenith observations using the so-called geometrical approximation (Hönninger et al., 2004) and second, from direct conversion of the zenith-sky observations using appropriate AMFs. For the present analysis, and considering that our test data were recorded under clear-sky aerosol-free conditions, zenith-sky HCHO AMFs were calculated using the UVspec/DISORT model (Mayer and Kylling, 2005; Hendrick et al., 2006) at the wavelength of 346 nm and for a typical HCHO profile peaking in the boundary layer. The HCHO content in the noon reference spectrum was derived using the geometrical approximation, so that both VCD evaluations (geometrical approximation and zenith-sky conversion) were constrained to agree at the time of the noon reference spectrum. The resulting time-series of retrieved HCHO VCDs are displayed in Fig. 9d, e, f for the same three polynomial settings. As can be seen only the third case, i.e. the DOAS evaluation using a 5th order polynomial, leads to consistent retrievals of HCHO VCDs using both geometrical approximation and zenith-sky conversion.

Although these results strengthen our confidence in the corresponding HCHO DSCDs, the question remains: what is causing the observed dependence on the polynomial order? The curvature of the zenith-sky daily variation observed when using polynomials of order 3 and 4 (black dots in the first two upper plots of Fig. 9) is striking and suggests an interference problem involving another absorber. As can be seen in Fig. 10, the Ring effect clearly displays a similar curved pattern. Additional tests

showed that none of the other parameters involved in the HCHO retrieval produces a similar shape.

The Ring effect (Grainger and Ring, 1962) is a well-known phenomenon responsible for filling-in of the solar and telluric lines in scattered sunlight spectra (e.g. Grainger and Ring, 1962; Wagner et al., 2009 and references therein). This effect is large in comparison to the faint absorption features of HCHO and it can therefore produce interferences if not well corrected in the DOAS evaluation. To investigate further, the sensitivity of our HCHO DSCDs to uncertainties in the Ring effect, additional test analyses were performed using different sources for the Ring cross-sections:

- Baseline case: cross-section calculated according to Chance and Spurr (1997)
- Case A: normalised cross-section generated according to Wagner et al. (2009)
- Case B: cross-section derived from SCIATRAN radiative transfer calculations in a Rayleigh atmosphere (A. Richter, personal communication, 2005)
- Case C: two cross-sections generated from Principal Component Analysis of a range of SCIATRAN calculations in an ozone containing atmosphere, according to Vountas et al. (1998)

Figure 11 shows the relation between changes in HCHO DSCDs and corresponding changes in the Ring fit coefficients when using different Ring cross-sections. One can see that HCHO dDSCD changes compared to the baseline ($\Delta \text{dDSCDs} = \text{dDSCD}(\text{new setting}) - \text{dDSCD}(\text{baseline})$) are linearly related to changes in the Ring fit coefficients. Comparing the retrieval cases using, respectively, a 3rd and a 5th order polynomials, it is also clear that the interference between HCHO and the Ring effect is much stronger when a 3rd order polynomial is used (Fig. 11a). This suggests that, for our analysis conditions, the use of a 3rd order polynomial introduces a misfit that activates the correlation between Ring and HCHO differential absorption features. In order to further test the stability of our retrievals with respect to this issue,

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a number of combinations of polynomials and Ring effect cross-sections were investigated and are summarized in Fig. 12. For each case, the root-mean square of the differences between HCHO VCDs retrieved using our two alternative VCD calculation methods are displayed. Overall, one can see that the sensitivity to the choice of the Ring effect cross-section is smallest when using a polynomial of order 5, for which more consistent results are obtained when changing the Ring datasets.

4.2 O₄ absorption cross-section

Another important interfering species in the HCHO fitting interval is the collisional dimer of molecular oxygen (O₄). Its absorption cross-sections are still poorly characterised due the difficulty of measuring them in laboratory under pressures and temperatures representative of atmospheric conditions. The Greenblatt et al. (1990) cross-sections are known to suffer from wavelength registration errors and have been measured under unsuitable high pressure conditions, not present in the ambient atmosphere. Alternatively, the Hermans et al. (2003) dataset provides cross-sections of overall better quality, however significant uncertainties still remain, particularly in the UV region below 360 nm. We tested both datasets, and came to the conclusion that significant interferences take place between O₄, HCHO and BrO absorption features in the 336.5–359 nm interval. Figure 13 shows the HCHO and BrO DSCDs, as retrieved from MAXDOAS measurements using the Hermans et al. and the Greenblatt et al. O₄ cross-sections. As can be seen, the Hermans dataset (our initial baseline for the intercomparison exercise) leads to larger HCHO columns but also to a larger spread in the BrO DSCDs retrieved at different viewing elevations, a feature that is not expected for a stratospheric absorber like BrO. In contrast, the BrO DSCDs derived using the Greenblatt O₄ cross-section appear to be more consistent. Similarly to the case of the polynomial discussed before, this suggests that a misfit to the O₄ absorption (larger in this case using the Hermans et al. dataset) activates a correlation between HCHO and BrO DSCDs. We will revisit the origin of these correlations in the next section. The linear relationship between HCHO, BrO and O₄ DSCDs changes when switching from the Hermans et al. to

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the Greenblatt et al. cross-sections, as is clearly apparent from Fig. 14. Based on these considerations, we conclude that the Greenblatt et al. O_4 cross-section so far remains the better choice for HCHO retrieval in the 336.5–359 nm region.

4.3 DOAS fitting interval

As already mentioned, the baseline HCHO fitting interval selected for the present study extends from 336.5 to 359 nm. This wavelength region, which includes three strong absorption bands of HCHO, has generally been recommended for HCHO DSCD retrievals. However, the absorption structures of HCHO and BrO are to some extent correlated in this wavelength interval, which has been identified as an issue for satellite BrO retrievals (Theys et al., 2011). Figure 15b graphically displays the correlation matrix of the different absorption cross-sections used in the HCHO fit. As can be seen, HCHO and BrO present the largest coefficient of correlation (around 0.55), which can be easily explained by the similarities of their differential absorption cross-sections (see Fig. 15a). In comparison, other species are less correlated. However, the coefficient of correlation between HCHO and O_3 and for other combinations involving O_3 , O_4 , BrO, NO_2 and Ring are not negligible. Such correlations may be expected to be dependent on the wavelength interval considered for the analysis. Therefore, in an attempt to identify the settings that would minimize the correlation matrix, calculations were performed for a range of fitting intervals starting between 332 and 338 nm and ending between 352 and 360 nm, in steps of 0.25 nm, in a similar way than what was done in Vogel et al. (2012). For each case, the root-mean-square of the non-diagonal elements of the correlation matrix was reported in Fig. 15c. Smaller correlations are clearly found for fitting intervals starting at short wavelengths. From visual inspection of Fig. 15c, one can conclude that the 333–358 nm wavelength range presents a local minimum of correlation, likely because of the addition of a BrO band at 334 nm in a region free of HCHO absorption. Note that this wavelength interval is similar to the one selected in Theys et al. (2011) for their satellite retrievals of BrO.

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To further explore the potential of this extended fitting interval on our HCHO MAX-DOAS retrievals, additional sensitivity tests were performed. Results again show large instabilities with respect to the Ring effect interference. This is illustrated in Fig. 16a, b, c where the HCHO DSCD retrieved in the 333–358 nm interval with a 5th order polynomial and O₄ Greenblatt et al. dataset, is displayed for different elevation angles and for different choices of the Ring cross-sections. As can be seen, the diurnal behaviour of the retrieved HCHO DSCDs has a large dependence on the source of the Ring cross-section used in the DOAS fit, and the corresponding HCHO VCDs calculated using the two methods introduced in Sect. 4.1 are generally inconsistent. These results suggest that the extended fitting interval that minimizes the BrO-HCHO interference is also more sensitive to Ring effect misfits. Therefore, any attempt to use this interval should be made with great care.

4.4 Recommended analysis settings

The sensitivity studies performed on BIRA data of 4 July 2009 revealed several possible optimisations of the HCHO DOAS retrieval, in order to minimise interferences and misfits related to polynomial order, Ring effect and O₄ and BrO absorption cross-sections. Compared to the settings used during the intercomparison exercise and presented in Table 2, the use of a 5th degree polynomial and the wavelength corrected O₄ Greenblatt et al. (1990) cross-section is recommended. Applied in the 336.5–359 nm wavelength interval, these changes reduce instabilities related to the Ring effect and lead to more consistent BrO DSCDs. Sensitivity tests involving other parameters revealed a comparatively small impact on the HCHO evaluations. This is further discussed in the next section where the systematic and random uncertainties on HCHO DSCD retrievals are reviewed in more details.

5 Error budget

The total uncertainties on the HCHO DSCDs retrieval can be divided into two categories: (1) the random errors mostly caused by measurement noise, and (2) errors affecting the slant columns in a systematic way.

5.1 Random uncertainties

Random errors in DOAS observations are primarily related to the measurement noise which, for silicon array detectors, is generally limited by the photon shot noise. Assuming that the errors of the individual detector pixels are uncorrelated and that the DOAS fit residuals are dominated by instrumental noise, the random contribution to the DSCD error can be derived from the DOAS least-squares fit error propagation (e.g. Stutz and Platt, 1996). Random errors are then adequately represented by the slant column fit errors. Any deviation with respect to these assumptions generally results in an overestimation of the random error, so one can consider to a first approximation that the DOAS DSCD error constitutes an upper limit of the true random error. For the instruments involved in this exercise, slant column errors varying between $\sim 1 \times 10^{15}$ molec cm^2 and $\sim 2 \times 10^{16}$ molec cm^2 were reported by the different groups as illustrated in Fig. 17a for 4 July 2009 at SZA less than 75° . As expected, scientific grade instruments (BIRA, Bremen, IUPH, WSU) display small errors of the order of 1×10^{15} molec cm^2 while mini-DOAS types of instruments (e.g. Mainz) are significantly noisier, typical errors reaching 5×10^{15} molec cm^2 or more. In order to better compare the actual performances of the different instruments, the DSCD errors were further normalised to a common integration time of 1 min. As can be seen in Fig. 17c, this largely improves the consistency between the scientific-grade instruments, which all display similar noise levels. The NASA and Mainz instruments, which use small and uncooled or less cooled detectors (see Table 1) have larger errors as is to be expected. More surprisingly, the Toronto data appear to display a similar level of noise, despite the fact that this instrument was using a large two-dimensional cooled array detector. This is likely related to

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5 a mechanical slit problem having occurred during the campaign that reduced the signal to 25 % of normal values. Also temperature instabilities have led to systematic features in the residuals that affected DSCD error estimates reported by this group. During part of the campaign, the Bremen instrument had a high frequency structure in the residuals
10 which could be accounted for by using an empirical correction. The problem could not be reproduced in the laboratory after the campaign and appears not to have impacted the NO₂ and HCHO retrievals. The INTA instrument (not on Fig. 17 because it was not measuring on 4 July) is showing similar level of noise. When normalised to the same integration time as the other instruments, JAMSTEC appears to have been the noisiest
15 system operated during the campaign.

5.2 Systematic uncertainties

Several important sources of systematic uncertainty have already been discussed in depth as part of our sensitivity analysis, which led us to propose optimised HCHO retrieval settings that minimise interference effects involving the polynomial closure term,
15 Ring effect, O₄, BrO and HCHO cross-sections. In this section, additional uncertainties are treated with the aim of providing a comprehensive error budget for dDSCDs, as it's the main parameter of interest for profiling inversion. This includes the impact of systematic uncertainties in absorption cross-sections as well as errors due to calibration uncertainties, in particular the slit function and the wavelength calibration, which are
20 key parameters for DOAS retrievals.

5.2.1 Absorption cross-sections

HCHO

Two sources of HCHO absorption cross-sections have been used in the literature, the Cantrell et al. (1990) spectrum and the Meller and Moortgat (2000) dataset. The
25 latter was adopted for our baseline. HITRAN recommends the use of the Cantrell dataset, re-scaled to correct for its systematically low bias (Chance and Orphal, 2011).

In the 336.5–359 nm interval and at the resolution of the BIRA spectrometer, the cross-sections differ by approximately 9%, a difference which was found to propagate directly to the slant column retrievals. The temperature dependence of the HCHO absorption cross-section is small, of the order of $0.05\% \text{ } ^\circ\text{K}^{-1}$ (De Smedt, 2011) and its effect was neglected here since the Meller and Moortgat cross-section was measured at 293°K , a temperature representative of ambient conditions during CINDI.

BrO

Two main sources of BrO cross-sections can be found in the literature: Wilmouth et al. (1999) and Fleischmann et al. (2004). These datasets are highly consistent in shape and their use was found to result in very small differences in the HCHO dDSCD, of the order of a few $10^{14} \text{ moleccm}^2$. For a median dDSCD of $3.8 \times 10^{16} \text{ moleccm}^2$ at 4° elevation, the difference is therefore less than 2%.

Ozone

The baseline intercomparison settings used the Bogumil et al. (2003) ozone absorption cross-sections. We have tested the impact of using the alternative dataset from Brion, Daumont and Malicet (BDM, Daumont et al., 1992; Malicet et al., 1995). The resulting HCHO dDSCDs were found to be larger by approximately $5 \times 10^{15} \text{ moleccm}^2$. For a median dDSCD of $3.8 \times 10^{16} \text{ moleccm}^2$ at 4° elevation, the difference is on the order of 13% during the day.

Ring effect

Although the cross-talk between HCHO and the Ring effect has been strongly reduced using the new baseline settings defined in Sect. 4 (see Fig. 10), some level of correlation persists between these parameters. As a result, HCHO uncertainties are expected to be linked to the strength of the Ring effect, which itself is a function of the geometry, SZA and aerosol content (Wagner et al., 2009). When considering the Ring cross-section that leads to the larger differences, (in most cases approach A), typical

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uncertainties on the HCHO DSCD reaches up to 5 to 12%.

NO₂

The baseline intercomparison settings used the Vandaele et al. (1996) NO₂ absorption cross-sections at 298° K. Switching to the alternative dataset of Burrows et al. (1998) HCHO DSCDs are found to vary by 2 to 5 %, depending on the NO₂ content.

O₄

The choice of the O₄ cross-section has been already largely discussed in Sect. 4.2. Adopting the Greenblatt et al. (1990) dataset, which minimises inconsistencies in the BrO slant columns, we further neglect residual errors related to O₄. It should be noted that in this study, a wavelength axis corrected version of the Greenblatt et al. dataset has been used (see e.g. Wagner et al., 2002), with a shift of -0.2 nm.

5.2.2 Instrumental slit function and wavelength calibration

Uncertainties in key instrumental calibration parameters may also be important. For example, imperfect characterisation of the slit function can lead to errors in the HCHO retrieval due to inappropriate convolution of the laboratory absorption cross-sections. This effect was tested by changing slightly the width of the measured slit function of the BIRA instrument (changes of ± 0.04 nm around the measured value of 0.38 nm Gaussian FWHM). This led, e.g. to changes in the HCHO dDSCD at 4° elevation of 1×10^{15} molec cm^{-2} , corresponding to relative differences of around 2%. In addition, small perturbations to the wavelength registration of the spectra were introduced in order to test the sensitivity of the retrievals to known uncertainties in the wavelength calibration procedure. Results indicate changes of the order of 2×10^{15} molec cm^{-2} for the 4° dDSCD (around 5%) for shifts of 0.02 nm. We therefore conclude that uncertainties in the instrumental slit function and the wavelength calibration have a minor impact.

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Another potentially important source of error already identified in the CINDI NO₂ study by Roscoe et al. (2010) is the accuracy of the pointing direction. Since this error source does not affect directly the DOAS evaluations, but more the determination of the corresponding air mass factors, we do not treat it explicitly here. It should of course be considered for a full error analysis of vertical column and/or profile concentration retrieval of HCHO. Another source of error to consider for the profile inversion is the atmospheric noise, due to the temporal and spatial variability of HCHO during the off-axis scan and the clouds interferences.

5.3 Overall error budget

Based on the results discussed above, an overall assessment of the total uncertainties on HCHO dDSCDs has been generated, including the main contributions of systematic and random errors, and is shown in Fig. 18. The figure summarises the main results from the sensitivity study for an elevation angle of 4° with a typical dDSCD of 3.8×10^{16} molec cm^2 , which is the median value during the 4 July 2009. For most cases, the retrieved HCHO dDSCDs fall within 15% of the values obtained with the optimised settings defined in Sect. 4. Assuming that the different effects are sufficiently uncorrelated with each other, we can sum all deviations in quadrature to obtain an estimate of the overall systematic uncertainty, which is represented by the black line in Fig. 18. On this basis, we estimate the total systematic uncertainties on HCHO dDSCDs to be of approximately 20% for measurements at 4° elevation, with a weak dependence on the SZA. Since some of the effects considered in our study are likely to be partly correlated, these values could be considered as upper limits, however, despite our efforts to include the most important sources of uncertainties in our sensitivity analysis, the need for possible additional terms cannot be excluded a-priori. Moreover, the results presented here are based on a single day of measurements and are influenced by the atmospheric conditions on that day. Therefore, arguably, the uncertainties reported here are to be interpreted as realistic conservative values.

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In Fig. 18, the random error is estimated for a typical integration time of one minute. We distinguish between two typical cases corresponding to low-noise scientific grade instruments and to mini-DOAS types of instruments, respectively. As can be seen for scientific grade instruments, the total errors on individual measurements are largely dominated by the systematic part. For mini-DOAS instruments, both random and systematic uncertainties contribute similarly. As already mentioned, the random uncertainty can be reduced by means of longer integration time and, for less sensitive mini-DOAS instruments, a trade-off between error and temporal resolution has to be made.

6 Conclusions

HCHO differential slant columns were retrieved from nine different MAXDOAS instruments jointly operated during the CINDI campaign in Cabauw, the Netherlands from June to July 2009. This exercise complements the formal semi-blind NO₂ and O₄ slant column intercomparison performed during the same campaign (Roscoe et al., 2010). To reduce the impact of uncertainties on retrieval parameters, common DOAS analysis settings were used by the different groups. In addition, 30-min averages were taken to reduce differences in temporal sampling. The HCHO differential slant columns retrieved by the different groups generally agree within 15 %, which is very satisfactory and almost as good as the agreement obtained for the NO₂ and O₄ intercomparison (Roscoe et al., 2010). This exercise shows that a large variety of MAXDOAS instruments of different grades and sensitivities can consistently measure HCHO columns within acceptable errors. However, while scientific grade instruments clearly demonstrate their ability to provide low-noise measurements at high temporal resolution (less than 30 min), less sensitive mini-DOAS instruments display significantly larger noise, which probably compromises their ability to deliver vertical profile information at the same temporal resolution.

A number of sensitivity tests was performed to investigate the sensitivity of HCHO retrievals to changes in DOAS analysis settings and input datasets. The study highlights

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the role of cross-correlation effects involving Ring effect, O_4 , BrO and HCHO absorption cross-sections and the order of the DOAS closure polynomial. Optimised retrieval settings are proposed with the aim to minimise correlation effects. Furthermore, systematic and random uncertainties are estimated for typical observation conditions. The largest systematic errors are found to be related to the Ring effect and to the uncertainties in HCHO and O_3 absorption cross-sections. We conclude that the overall systematic uncertainty on the HCHO DSCD retrievals is of the order of 20 % with a weak dependence on the solar zenith angle. Total errors are dominated by systematic effects for the scientific grade instruments, while both systematic and random uncertainties contribute at the same level for mini-DOAS-like instruments.

Finally, it must be pointed out that issues already identified during the semi-blind NO_2 and O_4 intercomparison (Roscoe et al., 2010) remain largely valid for the present study on HCHO. In particular, the accuracy of slant column measurements crucially depends (especially for low elevation angles) on the accuracy of the pointing direction, which should therefore be checked on a regular basis. Also, the temporal variability in tropospheric signals can be large, even if this variability is expected to be less important for HCHO than for NO_2 . For future MAXDOAS intercomparisons, a better synchronisation of the measurements should be considered to minimise the scatter possibly introduced by differences in measurement time.

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Table 1. Primary characteristics of the spectrometers taking part in the HCHO intercomparison campaign.

Institute	Measurement period for HCHO	FWHM (nm)	Sampling ratio (pixel/FWHM)	Detector characteristics	Integration Time	Reference
BIRA	13 Jun to 22 Jul	0.4	9	back-illuminated CCD, 2048 × 512 pixels (−30 °C)	60 s	Clémer et al. (2010)
INTA	7 Jul to 24 Jul	0.4	3.6	UV enhanced CCD, 1024 × 256 pixels (−40 °C)	90 s	Roscoe et al. (2010)
Bremen	8 Jun to 21 Jul	0.4	11.8	back-illuminated CCD, 2048 × 256 pixels (−35 °C)	40 s	Wittrock et al. (2004)
Heidelberg	17 Jun to 24 Jul	0.5	6.5	back-illuminated CCD, 2048 × 256 pixels (−30 °C)	60 s	Frieß et al. (2011)
JAMSTEC	8 Jun to 24 Jul	0.7	8	uncooled CCD, 3648 pixels	5 min	Irie et al. (2011)
NASA	22 Jun to 20 Jul	0.6	5	uncooled CCD, 2048 × 14 pixels	16 s	Herman et al. (2009)
WSU	21 Jun to 5 Jul	0.83	8	back-illuminated CCD, 2048 × 512 pixels (−70 °C)	Typical 1.2 s	Herman et al. (2009)
Toronto	30 Jun to 4 Jul	0.2–0.8	7–12	back-illuminated CCD, 2048 × 512 pixels (−72 °C)	~ 2 min	Fraser et al. (2009)
Mainz	21 Jun to 10 Jul	0.6	8	Stabilised CCD, 2048 pixels (4 °C)	60 s	Shaiganfar et al. (2011)

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Table 2. Baseline DOAS analysis settings used for HCHO slant column retrieval during the intercomparison exercise.

Parameter	Specification
Fitting interval	336.5–359 nm
Wavelength calibration	Calibration based on reference solar atlas (Chance and Kurucz, 2010)
Cross sections	
HCHO	Meller and Moortgat (2000), 293° K
O ₃	Bogumil et al. (2003), 223° and 243° K, <i>I</i> ₀ -corrected
NO ₂	Vandaele et al. (1996), 298° K, <i>I</i> ₀ -corrected
BrO	Fleischmann et al. (2004), 223° K
O ₄	Hermans et al. (2003) (available at: http://spectrolab.aeronomie.be/o2.htm)
Ring effect	Chance and Spurr (1997)
Closure term	Polynomial of order 3 (corresponding to 4 coefficients)
Intensity offset	Linear correction
Wavelength adjustment	All spectra shifted and stretched against reference spectrum

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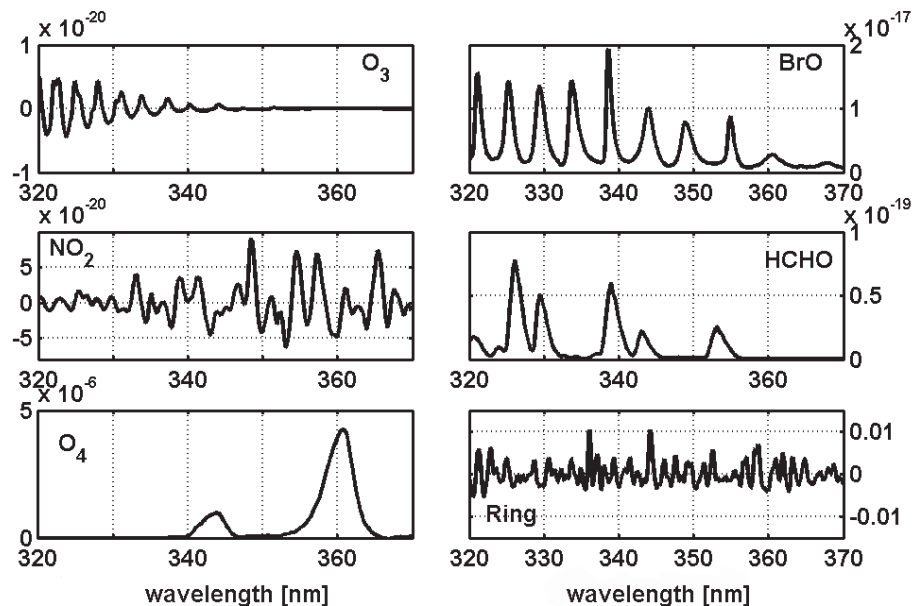


Fig. 1. Absorption cross-sections of the main trace gases absorbing in the 320–370 nm spectral range and of the Ring effect. All absorption cross-sections are expressed in $\text{cm}^2 \text{molec}^{-1}$, except O_4 ($\times 10^{-40} \text{cm}^5 \text{molec}^{-2}$) and have been smoothed to match the resolution of the BIRA-IASB instrument (0.38 nm FWHM).

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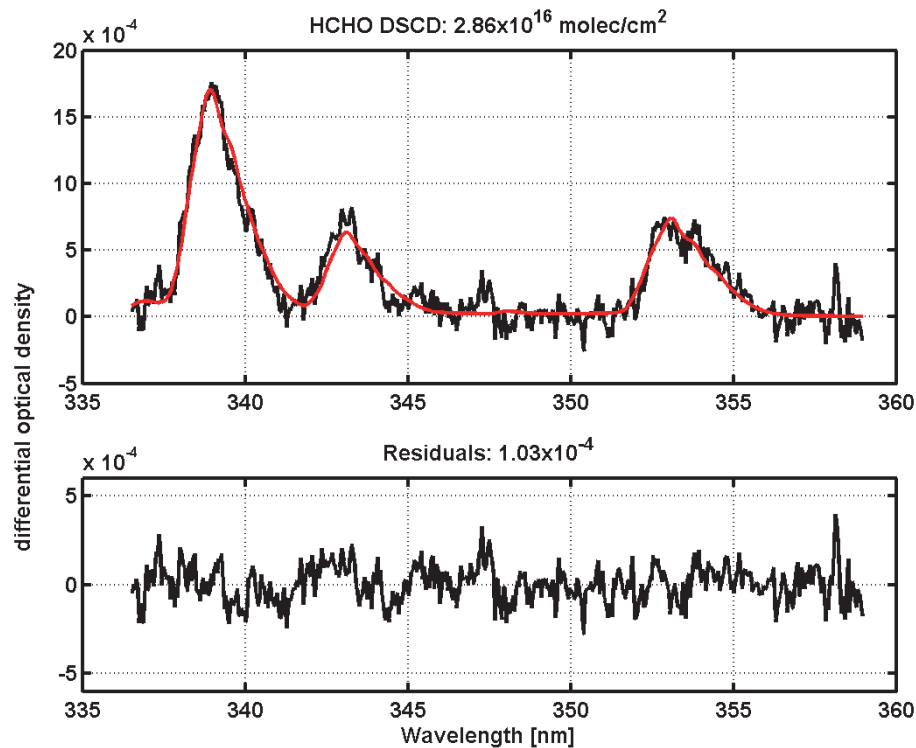


Fig. 2. Example of a HCHO slant column fit obtained with the BIRA-IASB instrument on 30 June 2009, around 14:30 UT, at 4° elevation angle and 43° solar zenith angle.

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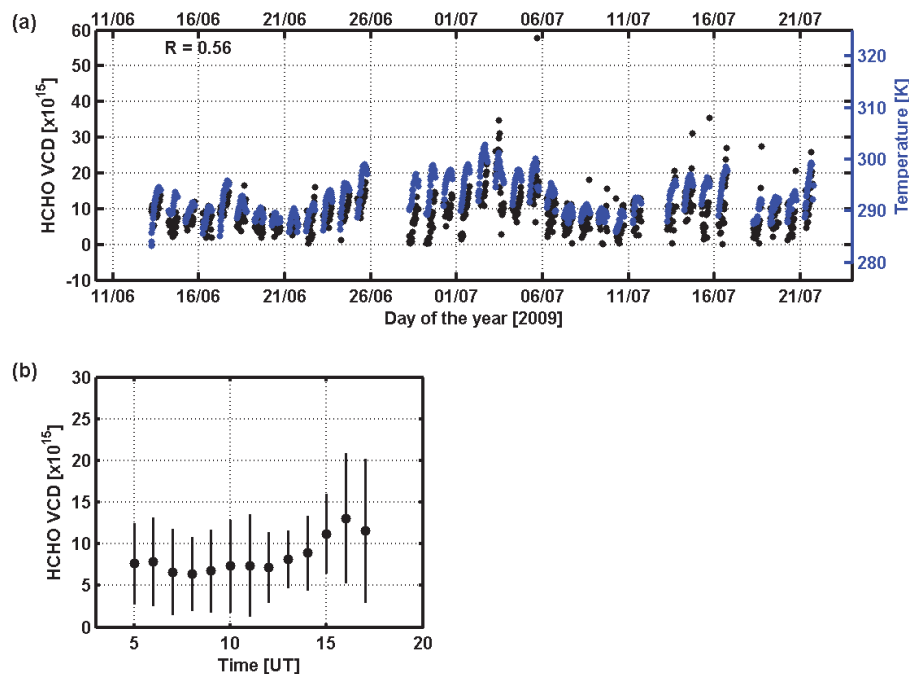


Fig. 3. HCHO vertical columns (molec cm^{-2}) retrieved from BIRA-IASB MAXDOAS measurements at 30° elevation using a geometrical approximation. **(a)** Time series of observations performed during June and July 2009. Ambient temperature measurements are superimposed in blue. **(b)** Corresponding mean diurnal variation of the HCHO vertical column. Error bars denote the one sigma standard deviations of individual measurements around the mean values.

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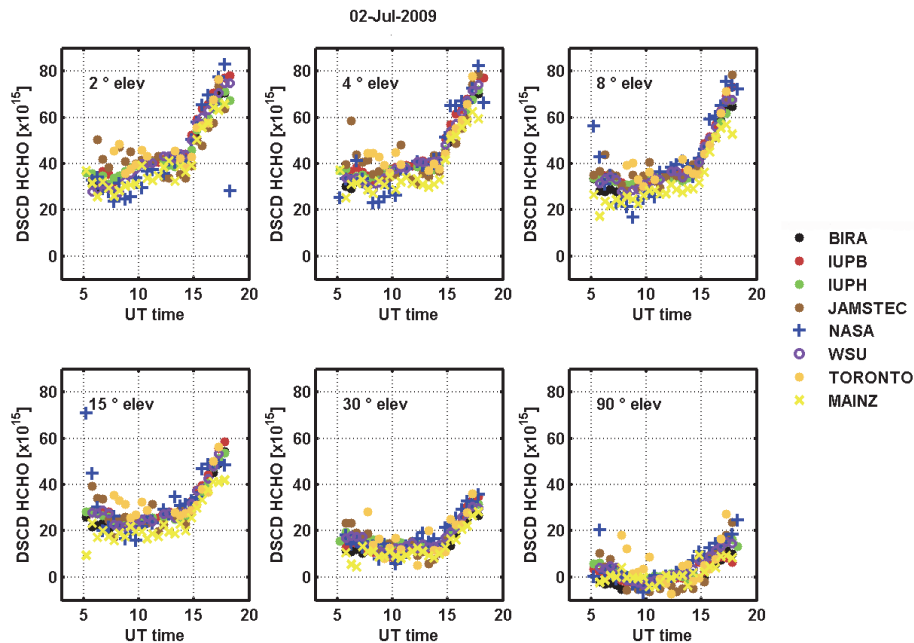


Fig. 4. Diurnal evolution of the HCHO DSCD measured on 2 July 2009 and averaged in 30-min bins, for the different instruments involved in the intercomparison. Units are molec cm^{-2} .

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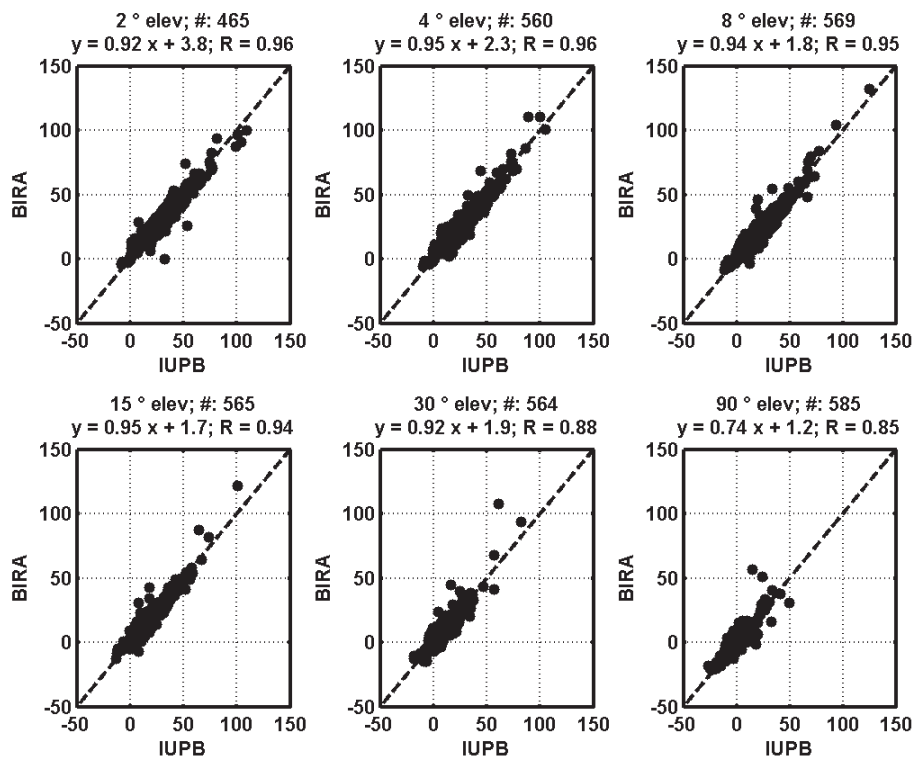


Fig. 5. Scatter plots of the HCHO DSCD measured by BIRA against those measured by IUP Bremen. Data from the whole campaign were averaged in 30-min bins and are expressed in units of 10^{15} molec cm^{-2} . Statistical parameters derived from each regression plot (number of points #, linear regression equation and correlation coefficients R) are given for each elevation above the corresponding subplots.

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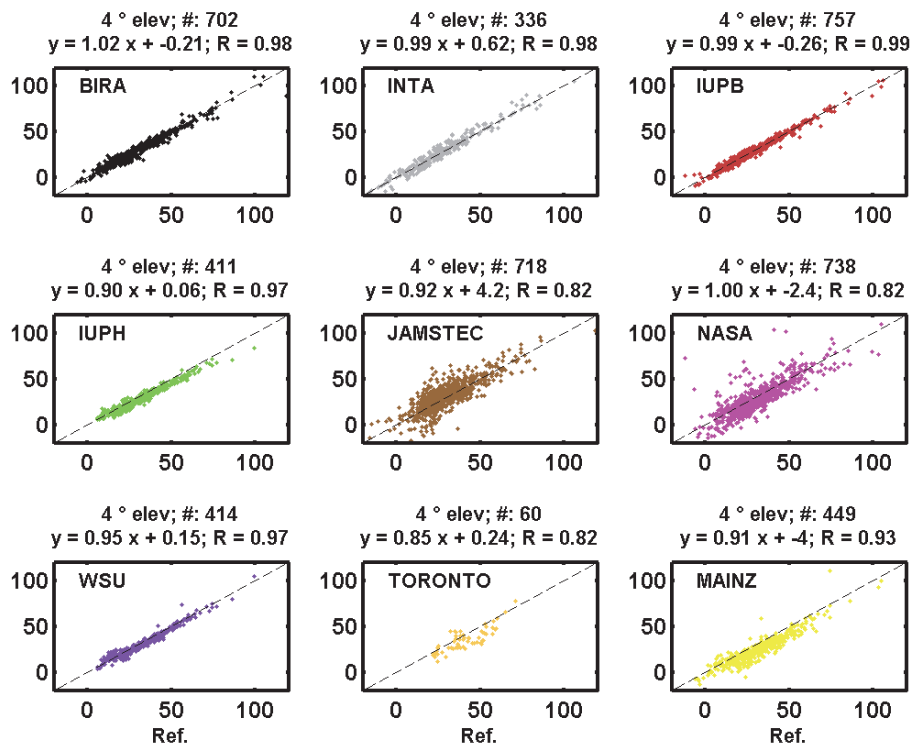


Fig. 6. Scatter plots of HCHO DSCDs (molec cm^{-2}) measured by each instrument compared to the campaign reference dataset (see text), for the case of measurements at the 4° elevation angle. Statistical parameters derived from the regressions (number of points, linear regression equation and correlation coefficients) are given on top of each subplot.

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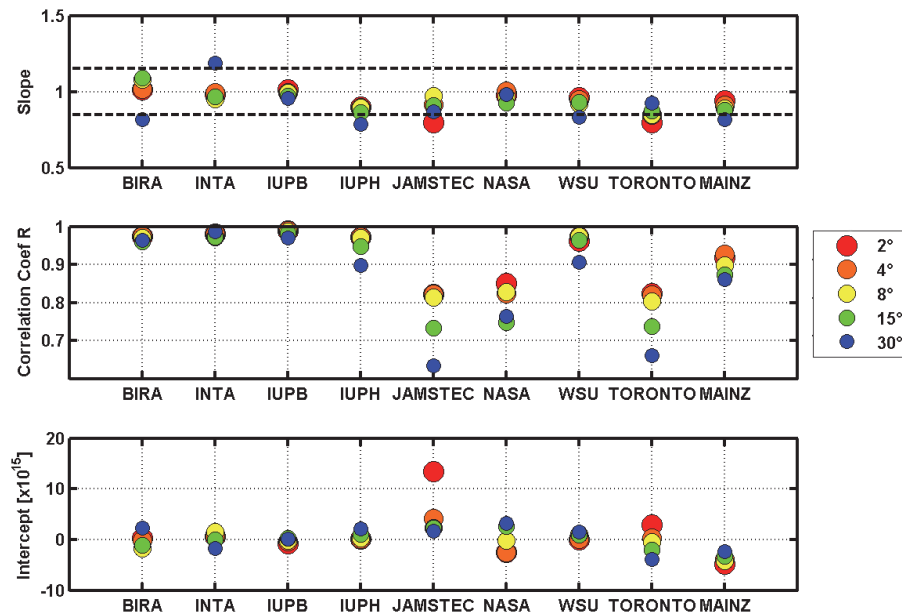


Fig. 7. Straight-line slopes, correlation coefficients and intercepts of HCHO slant columns against those of the reference, for each instrument and all off-axis elevation angles. The dotted lines in the first panel correspond to values of 1.15 and 0.85.

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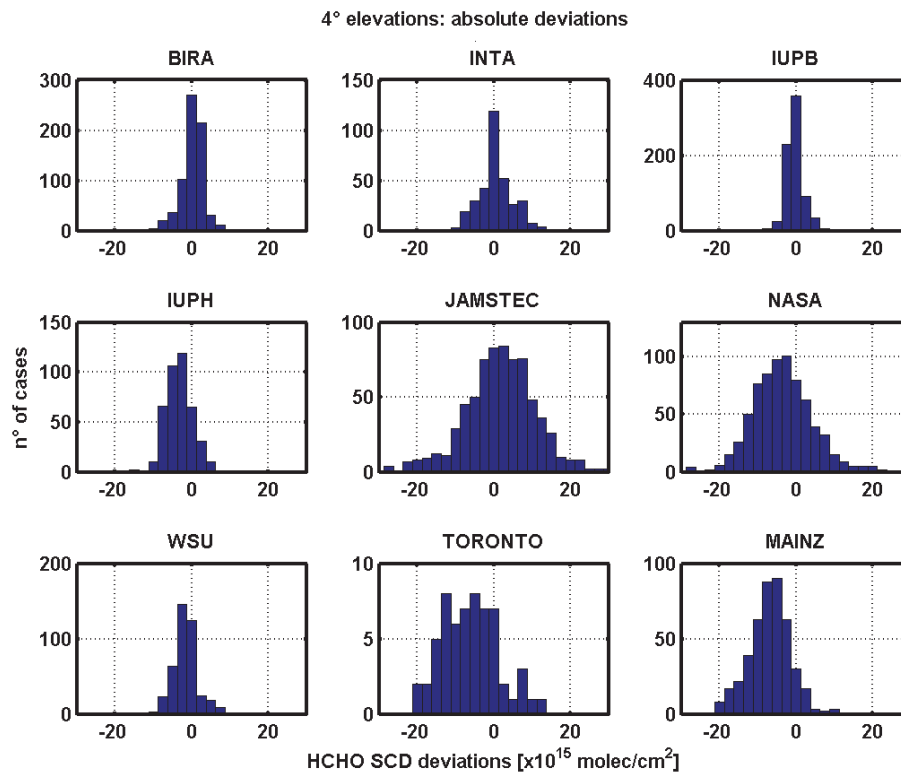


Fig. 8. Histograms of the HCHO DSCD absolute deviations (10^{15} molec cm⁻²) after straight line fits of each instrument's data to that of the reference set, for the case of measurements at the 4° elevation angle, and for the whole campaign.

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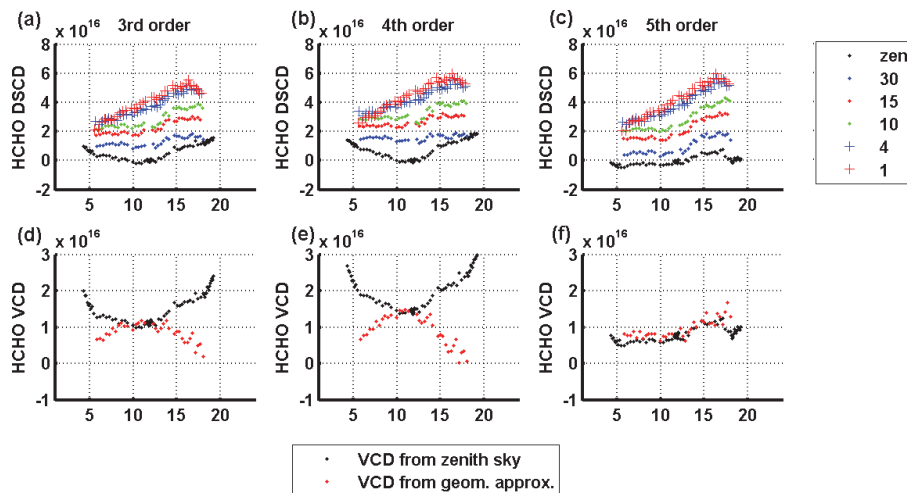


Fig. 9. Effect of the choice of the polynomial order used in the DOAS fit. **(a), (b), (c)**: impact on the HCHO DSCDs for the different elevation angles; **(d), (e), (f)** impact on the corresponding HCHO VCDs obtained using two different methods described in the text.

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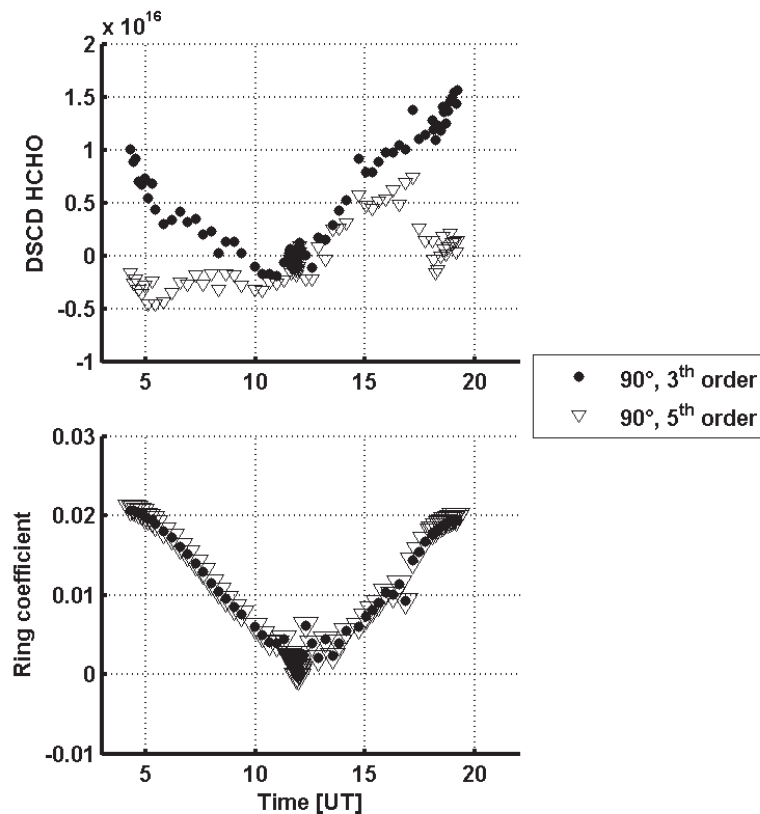


Fig. 10. Daily variation of the zenith HCHO DSCD columns and of the Ring coefficients for retrievals with a 3rd and 5th order polynomial. A strong correlation is observed for HCHO and Ring when a 3rd order polynomial is used in the DOAS retrieval.

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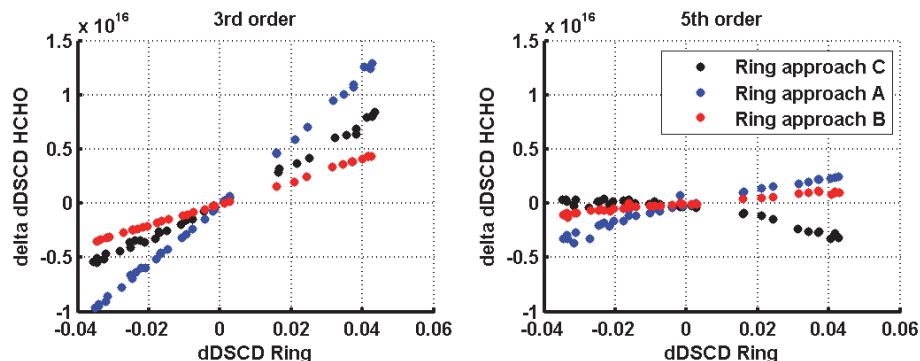


Fig. 11. Sensitivity of HCHO dDSCD to changes in the Ring cross-section used in the DOAS fitting procedure, expressed as the difference to the baseline scenario, for data recorded on 4th July 2009. The two panels present the results for different orders of the polynomial used in the DOAS fit.

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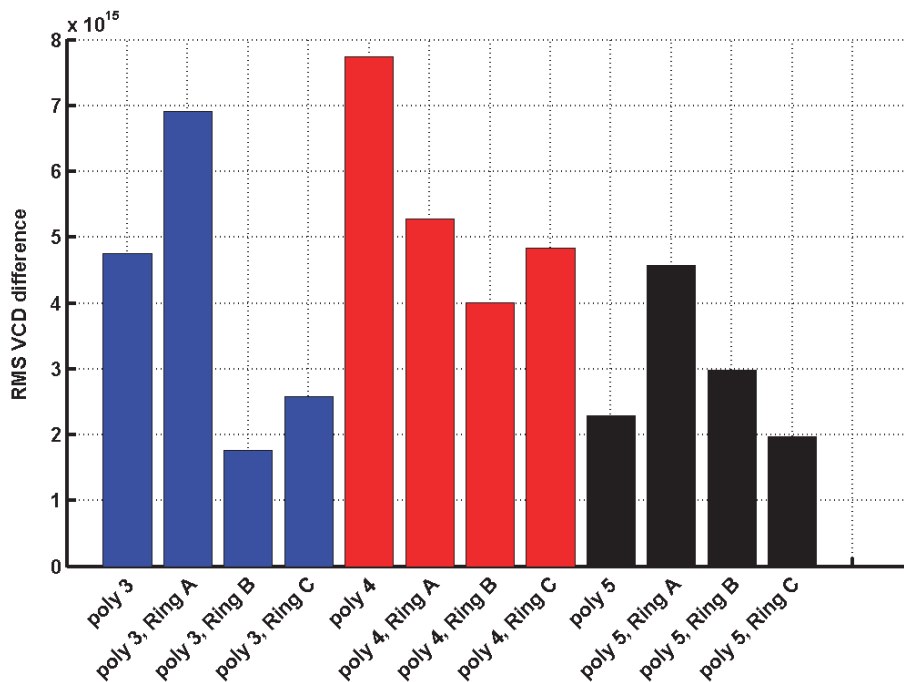


Fig. 12. Root-mean-square of HCHO VCD differences (molec cm^{-2}) obtained using two alternative methods for the calculation of vertical columns (see text and Fig. 9). Various choices of retrieval settings are tested using polynomials of degree 3 (blue), 4 (red) and 5 (black) respectively and different Ring effect cross-sections. Optimal stability (corresponding to smallest HCHO VCD differences and larger coherence between the results with different Ring choices) is obtained for cases using a polynomial of degree 5.

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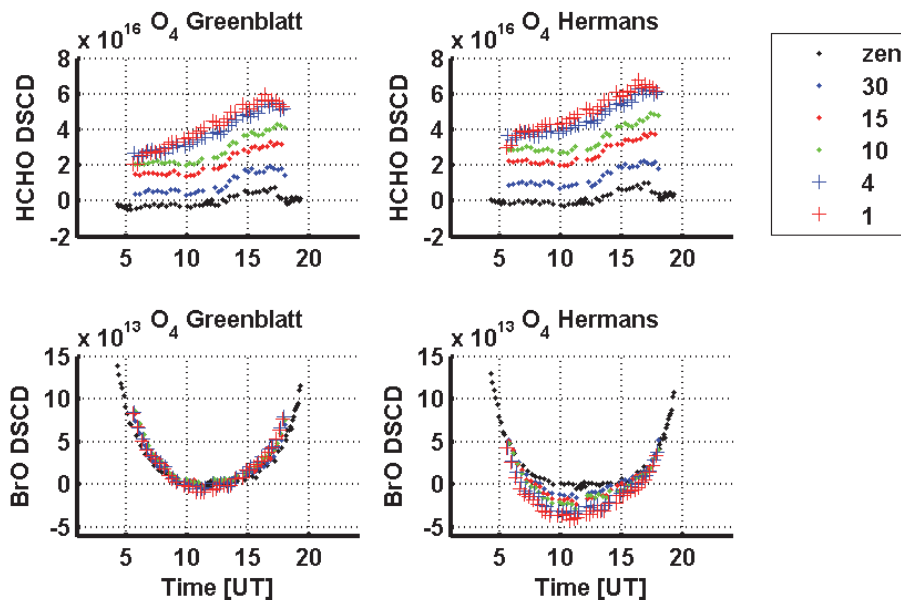


Fig. 13. Effect of the choice of the O₄ cross-section used in the DOAS fit, on HCHO and BrO DSCDs columns (molec cm⁻²). Results (a) and (c) are obtained with the Greenblatt et al. (1990) cross-sections while (b) and (d) are obtained using the Hermans et al. (2003) cross-sections.

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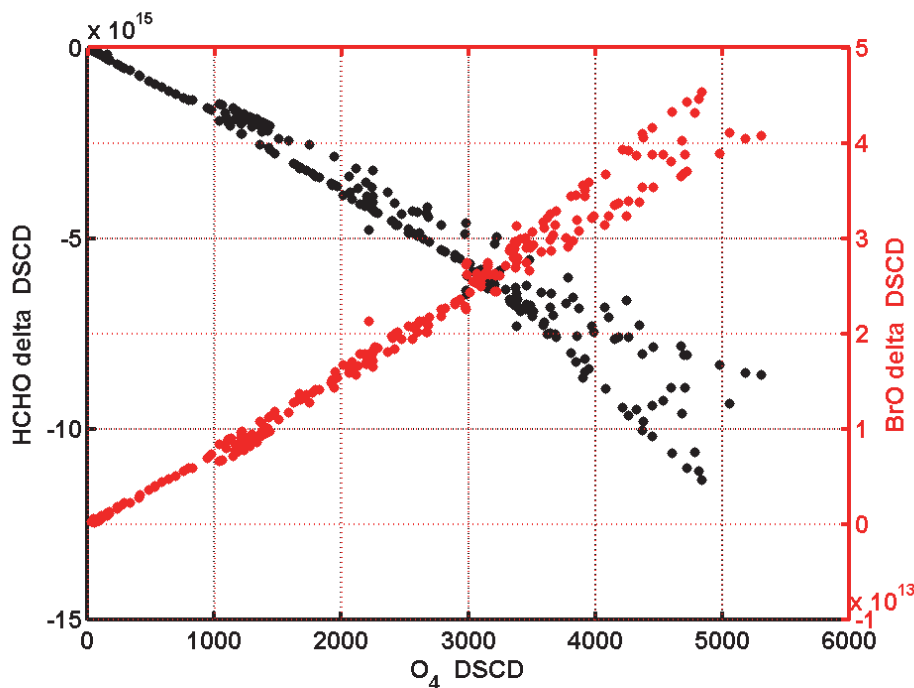


Fig. 14. Change in retrieved HCHO (black) and BrO DSCDs (red) (delta dDSCDs = dDSCD(new setting) – dDSCD(baseline); in molec cm^{-2}) when exchanging the Hermans et al. (2003) O₄ absorption cross-section for the Greenblatt et al. (1990) dataset, expressed as a function of the O₄ DSCD values (in 10^{40} molec 2 cm $^{-5}$).

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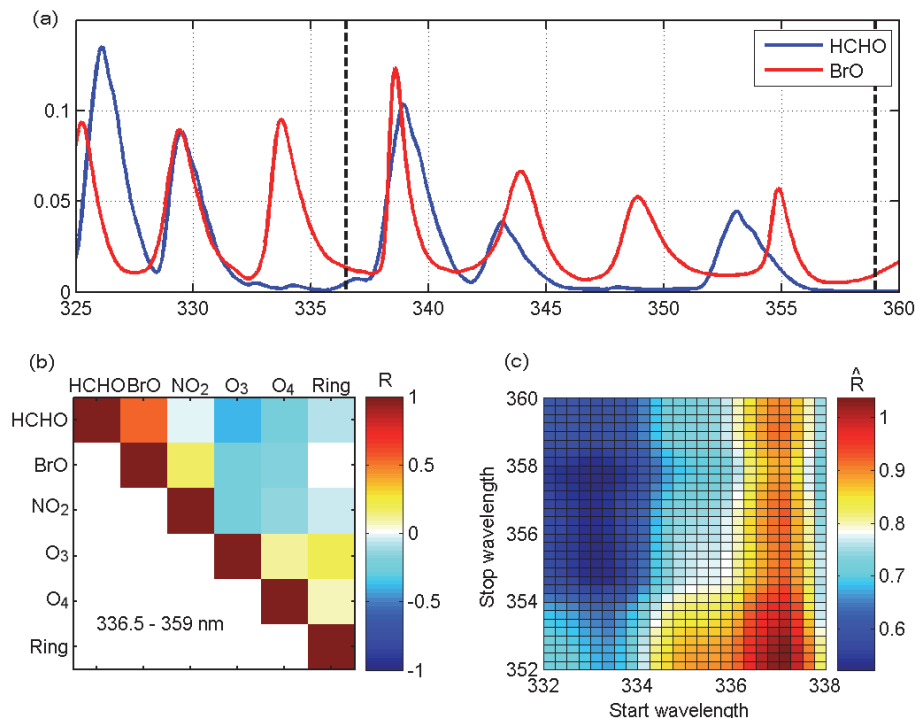


Fig. 15. (a) HCHO and BrO absorption cross-sections in the 325–360 nm wavelength range convolved at the resolution of the BIRA instrument (0.38 nm FWHM) and normalized in arbitrary units. (b) Correlation matrix of the absorption cross-sections used for HCHO DOAS retrievals in the 336.5–359 nm interval. (c) Overall correlation (expressed as the root-mean-square of the non-diagonal elements of the correlation matrix) for different wavelength intervals in the 332–360 nm wavelength range.

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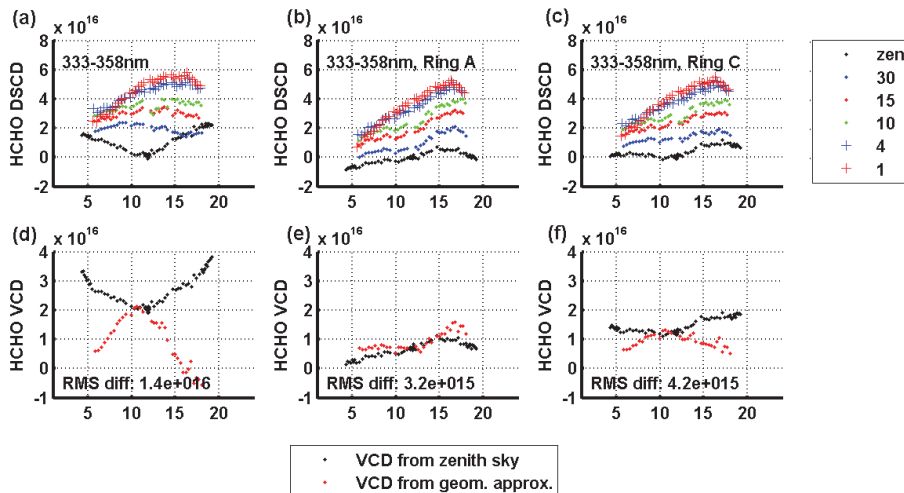


Fig. 16. Same as Fig. 9, but for an analysis in the 333–358 nm wavelength region, with a 5th order polynomial and the O₄ Greenblatt et al. dataset, and for different choices of the Ring cross-sections.

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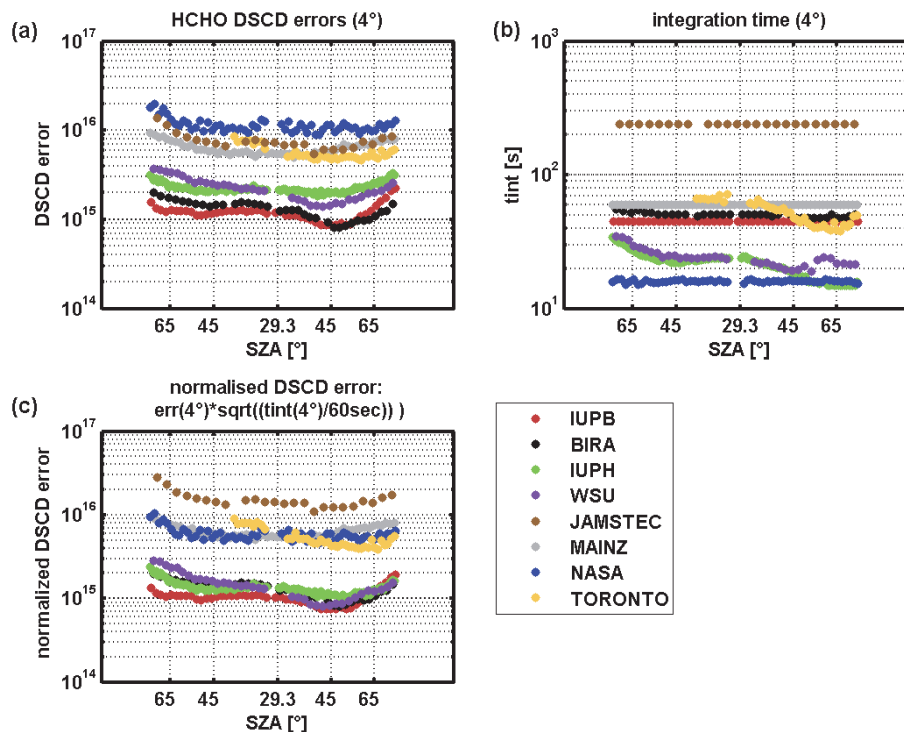


Fig. 17. Comparison of HCHO DSCD errors (molec cm^{-2}) retrieved by each participating group for the case of 4° elevation, based on measurements from 4 July 2009. **(a)** DSCD errors from DOAS evaluations, **(b)** corresponding integration times, **(c)** DSCD errors normalized by their integration times.

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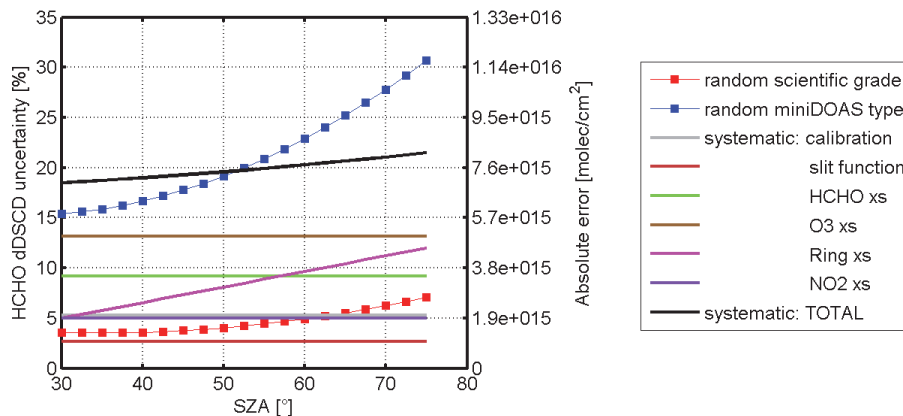


Fig. 18. Summary assessment of the error budget on HCHO dDSCD at 4° elevation, as a function of the SZA. Random uncertainties are typical of low-noise scientific grade instruments (red dots) and of mini-DOAS types of instruments (blue dots) for a typical integration time of 1 min.

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