Long-term trends of NO, above northern mid-latitudes as inferred from Jungfraujoch, HALOE and ACE-FTS solar observations.

INTRODUCTION

The NO_y family of gases, defined as NO + NO₂ + NO₃ + $2 \times N_2O_5$ + HNO₃ + HNO₄ + CIONO₂ + BrONO₂, plays an important role in the ozone depletion (NO_x catalytic cycle Crutzen 1970). At the Jungfraujoch observatory, FTIF spectrometers measure since 1984 the four most abundant members of NO_y i.e. NO, NO₂, HNO₃ and CIONO₂. Their sum is a good proxy of NO, and will be noted here NO, (the mos important missing gas being N_2O_5).

FTIR DATA SET

- Infrared solar absorption spectra recorded since 1984 at the Jungfraujoch observatory (Swiss Alps, 3580 m a.s.l.)
- > 2 FTIR spectrometers (home-made and Bruker 1 20 HR) ▶ high resolution (82 to 175 cm max. OPD)
- ► only clear-sky conditions
- ANALYSIS ► total columns
- > monthly means (avoid higher weight for period with many
- observations) > SFIT 1 and 2
- ► P, T profiles from NCEP
- spectral micro-windows
- NO: 1899.85-1900.20, 1902.92-1903.36 and 1912.70-1912.86 cm
- NO₂: 2914.51-2914.86 cm⁻¹
- HNO3: 868.75-869.75 cm
- CIONO₂: 779.3-780.6 then 780.050-780.355 cm⁻¹
- N2O: 2481.3-2482.6, 2526.4-2528.2, 2537.85-2538.8
- and 2540, 1-2540, 7 cm⁻¹ ► NO and NO₂: empirical correction for diurnal variation

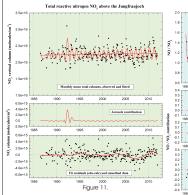
TRENDS DERIVATION

- ▶ multiple regression model, including a linear trend, a seasonal component and anomalies from various atmospheric parameters (Bodeker et al, 1998):
- solar flux (10.7 cm wavelength, measured at Ottawa / Penticton, Canada) stratospheric aerosol optical depth (15-35 km, 20° N -
- 50° N) (Vernier et al. 2011)
- tropopause height, calculated from NCEP P,T profiles • other investigated parameters: NAO (North Atlantic Oscillation), QBO (Quasi-Biennal Oscillation), pressure,
- stratospheric temperature only statistically significant parameters have been kept

Figures 6 to 9. Upper frames: partial columns of NO, NO₂, HNO₃, CIONO₂ derived from the HALOE (blue dots) and ACE-FTS (black dots) satellite experiments. The red curves correspond to the best fits to the daily means with a linear trend and a 6-term Fourier series, to characterize the long-term changes

and seasonal modulations of our target species. HALOE partial columns have been calculated above 30 mbar pressure level and for latitude belt from 42° N to 52° N. ACE partial columns are for latitude belt from 41°N to 51°N, and for altitude range of 28 - 55 km, 20 - 42 km, 20 - 37 km and 20 - 31 km, respectively for NO, NO $_2$, HINO $_3$ and CIONO $_2$. Lower frames show the residuals of the fit (measured - model) (black dots)

together with a smoothed curve (red lines) of these residuals. The minima of NO and NO, in 2007 are clearly visible in ACE-FTS data.



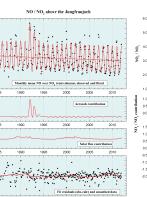
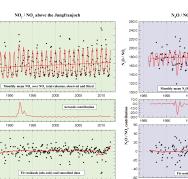
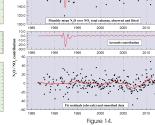


Figure 11





/	gas	instrum.	begin	end	ref. year	trend (%/yr)9	5 % sigma	method
I	NO	FTS	1984.8	2012.0	1984.0	-0.25	0.12	model & LS
		FTS	1991.8	2004.2	1991.8	0.20	0.28	bootstrap
		HALOE	1991.8	2004.2	1991.8	0.77	0.20	bootstrap
I		FTS	2004.1	2011.7	2004.1	-0.73	0.66	bootstrap
L		ACE	2004.1	2011.7	2004.1	-0.99	0.83	bootstrap
l	NO ₂	FTS	1985.2	2012.0	1984.0	0.08	0.12	model & LS
l		FTS	1990.0	2010.0	1990.0	-0.36	0.09	model & LS
I		SAOZ	1990.0	2010.0	1990.0	-0.37	0.11	model & LS
		FTS	2004.1	2011.7	2004.1	0.20	0.57	bootstrap
L		ACE	2004.1	2011.7	2004.1	0.14	1.42	bootstrap
	HNO ₃	FTS	1985.8	2012.0	1984.0	0.31	0.20	model & LS
		FTS	2004.1	2011.7	2004.1	2.08	0.94	bootstrap
L		ACE	2004.1	2011.7	2004.1	0.54	0.27	bootstrap
Γ	CIONO ₂	FTS	1986.5	1995.0	1995.0	3.57	1.01	model & LS
		FTS	1995.0	2012.0	1995.0	-0.73	0.34	model & LS
		FTS	2004.2	2011.7	2004.2	0.82	1.07	bootstrap
L		ACE	2004.2	2011.7	2004.2	0.50	0.50	bootstrap
	NOx	FTS	1985.1	2012.0	1984.0	0.08	0.16	model & LS
I	NO _y	FTS	1986.5	2012.0	1984.0	0.11	0.13	model & LS
Γ	N ₂ O	FTS	1984.4	1994.0	1994.0	0.46	0.06	model & LS
		FTS	1994.1	2012.0	1994.0	0.25	0.03	model & LS
ľ	NO/NO ₂	FTS	1985.2	2012.0	1984.0	-0.29	0.12	model & LS
1		FTS	1986.5	2012.0	1984.0	0.53	0.23	model & LS

FTS 1986.5 2012.0 1984.0 0.16

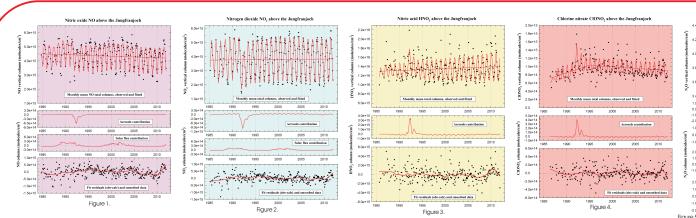
Canadian satellite SCISAT-1.

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- GAW-CH program (MeteoSwiss, Zürich); HFS Jungfraujoch A3C and AGACC-II (PRODEX and SSD programs from BELSPO, Brussels) ACE-FTS team, Canadian Space Agency (CSA) and NSERC, Canado

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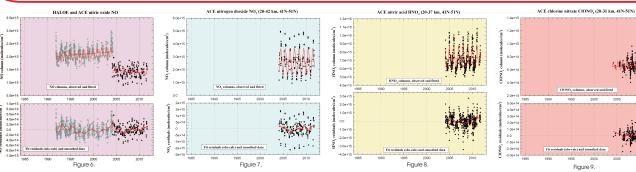
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Figures 1 to 5. Upper frames: monthly mean total columns derived from Jungfraujoch FTIR spectra (black dots), together with best fit of the regression model and linear trends (red lines), respectively for NO, NO₂, HNO₃, CIONO₂ and NgO. For CIONO₂, two linear trends have been included in the model, to take into account the decrease of this gas after 1995, as a consequence of the limitation of CI emissions (Montreal protocol). Two linear trends have also been used for N.O. Middle frames give the contribution of aerosol optical depth (AOD), solar flux, NAO and tropopause (when applicable) to the total columns. The large aerosol perturbation from the Mt. Pinatubo eruption on June 15, 1991, is clearly visible, sequestering NO, into HNO,. The eleven year solar cycle influences NO and NO, total columns by a few percents. A summary of these contributions can be found in Table 1.

Lower frames show the residuals of the fit (measured - model) (black dots) together with a smoothed curve (red lines) of these residuals. Note the decrease of NO, NO, and HNO, during 2006-2007, well visible in the smoothed residuals, so well in Jungfraujoch FTS data and in ACE data. This decrease is followed by an increase in 2008, then by a new ease in 2009 and again by a large increase in 2010.





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good agreement, except for HNO,

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Figures to 11 to 14: same as Fig. 1 to 5, for NO $_{\!\gamma}^{\,*}$ and for the ratios NO over NO., NO.^{*} over NO., and N.O over NO. NO_{y}^{*} shows no significant trend, but the minimum of 2007-2009 and the

increase in 2010 are well marked. Although N_2O is the source of NO_v^* , NO_v^* is not increasing at the same rate (0.31 \pm 0.02 %/year for N,O, 0.11 \pm 0.13 %/year for NO, see Table 2). This difference is due to increasing CO_ concentrations cooling the stratosphere (Rosenfield and Douglass, 1998) and to ozone and halogens changes in the stratosphere (McLinden *et al*, 2001). NO, is increasing at a rate of 0.53 \pm 0.23 %/year, mainly due to the

positive trend of HNO₃, the most abundant NO₂ species.

NO/NO₂ is decreasing at a rate of -0.29 ± 0.12 %/year. This decrease is due to increased chlorine loading in the atmosphere, which increases the rate of the reaction (NO + CIO \rightarrow NO, + CII. The increase of NO/NO, is expected in the 21st century, as a result of the decreasing chlorine loading and of CO₂-induced stratospheric cooling, which slows the temperaturelependent reaction [NO + $O_3 \rightarrow NO_2 + O_2$] (Revell et al., 2012).

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